

## **Diabatic-Dynamical Interaction in the General Circulation (lecture 1)**

In the coming period we are going to study the **interaction between dynamics and diabatic processes**.

### **Diabatic processes:**

*Absorption and emission of radiation*

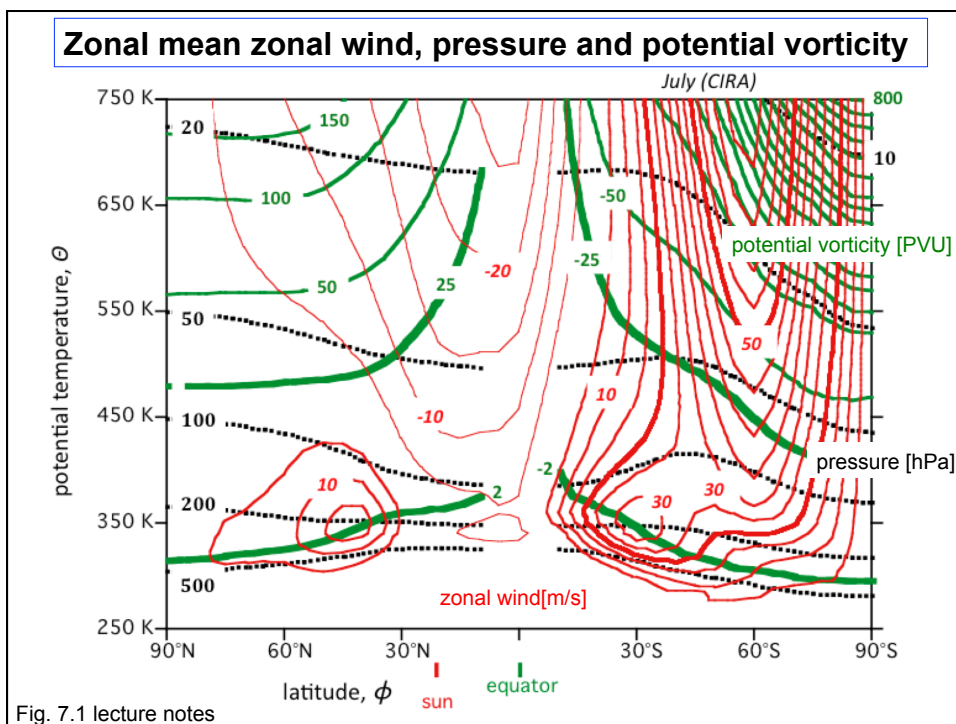
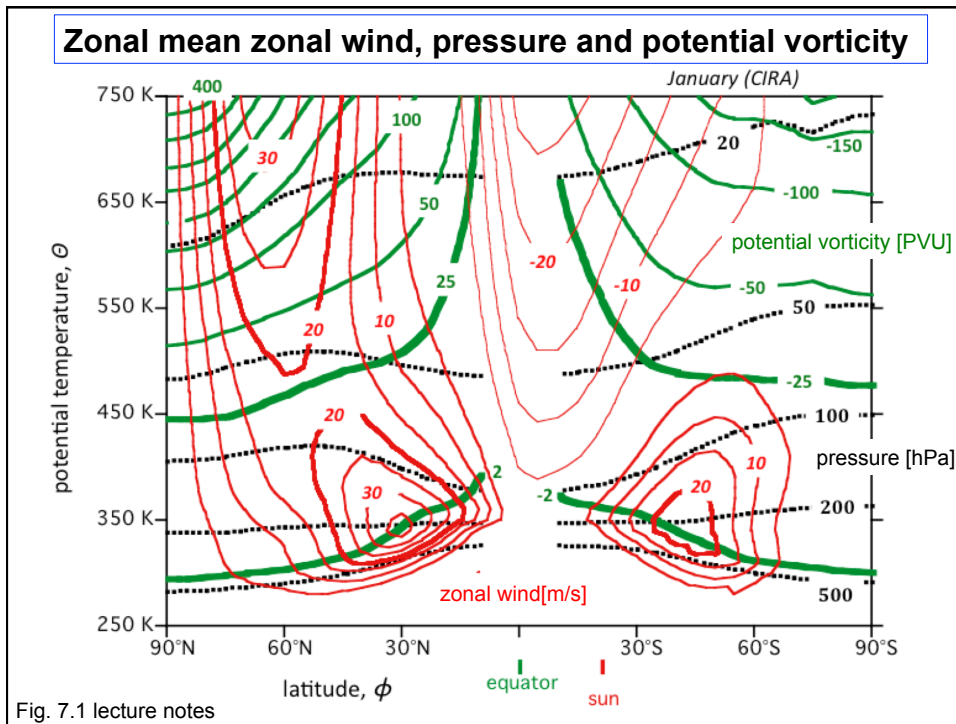
*Energy transformations associated with the water cycle*

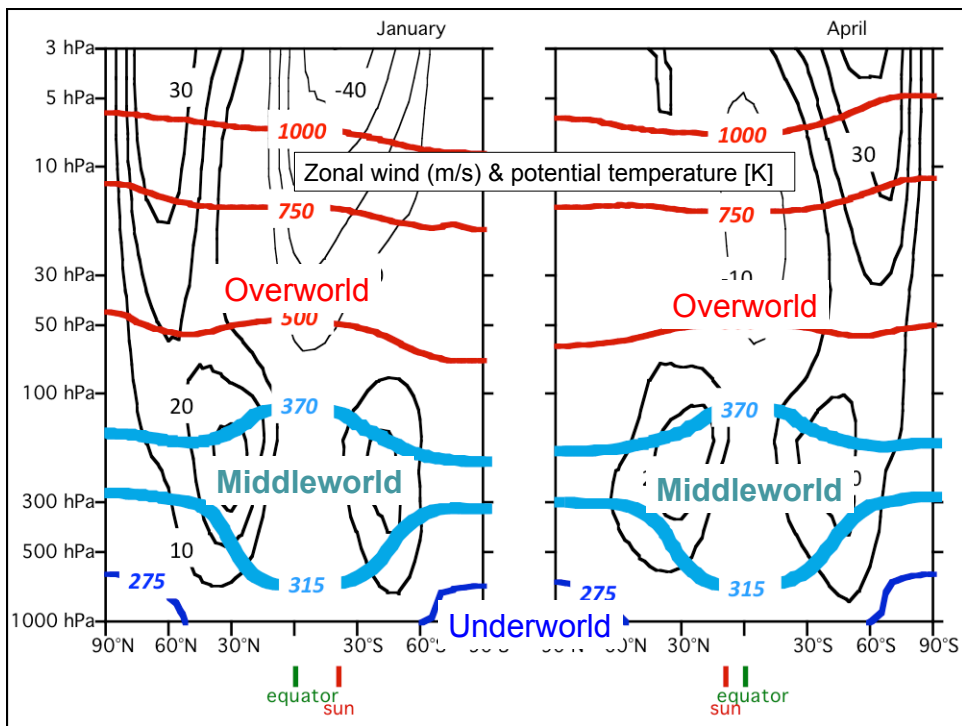
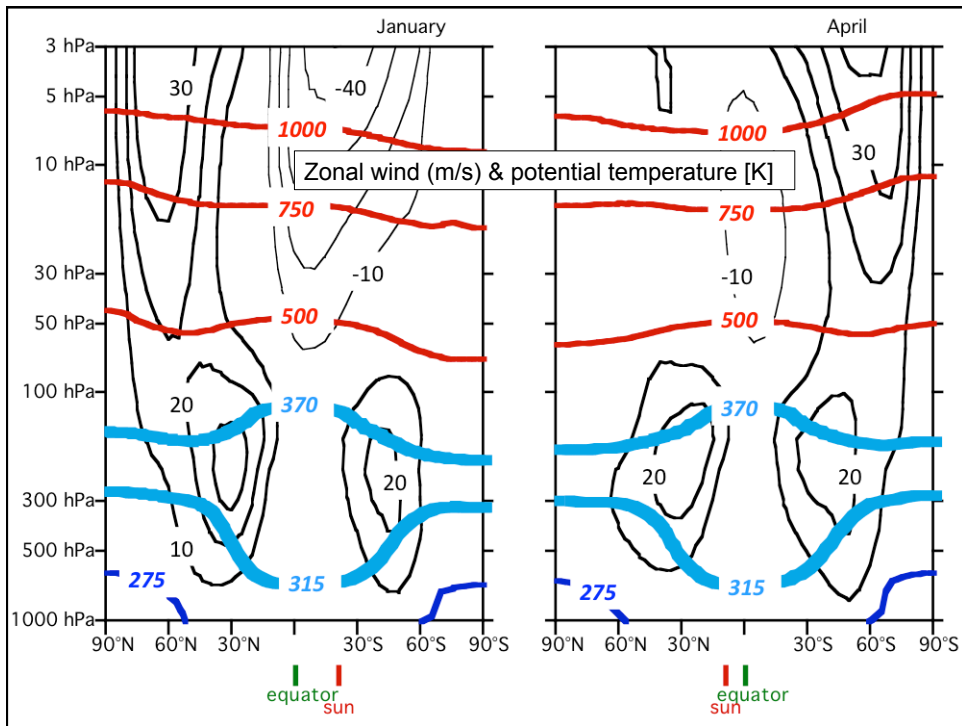
### **Dynamics:**

*Atmospheric circulation*

<http://www.staff.science.uu.nl/~delde102/C&HC.htm>

## The General Circulation





## What do we want to understand?

Position and intensity of subtropical jet

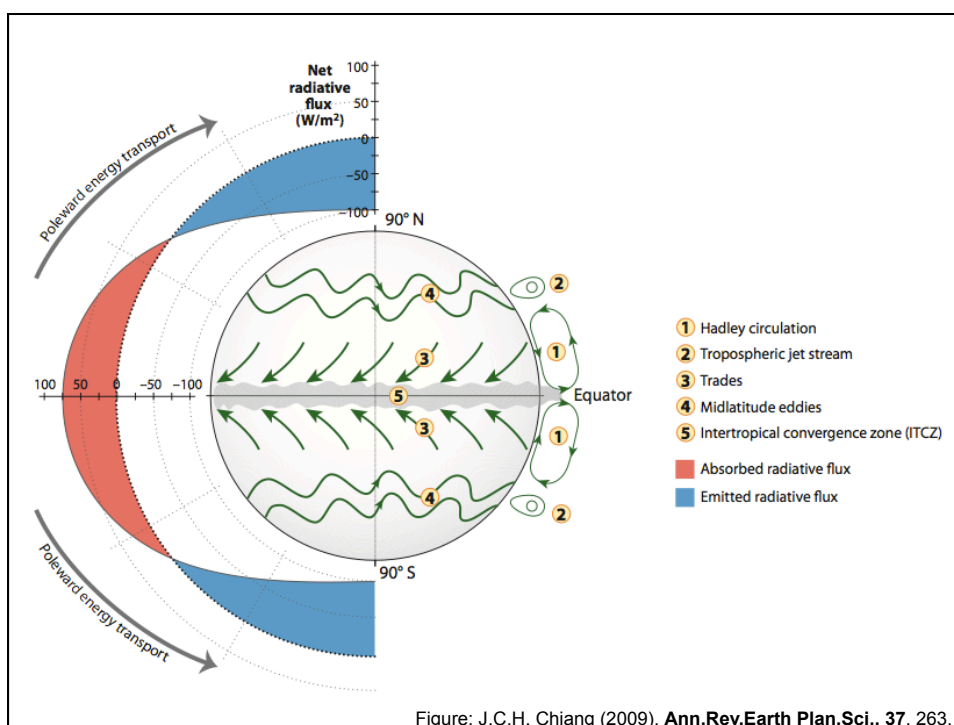
Cold tropical tropopause

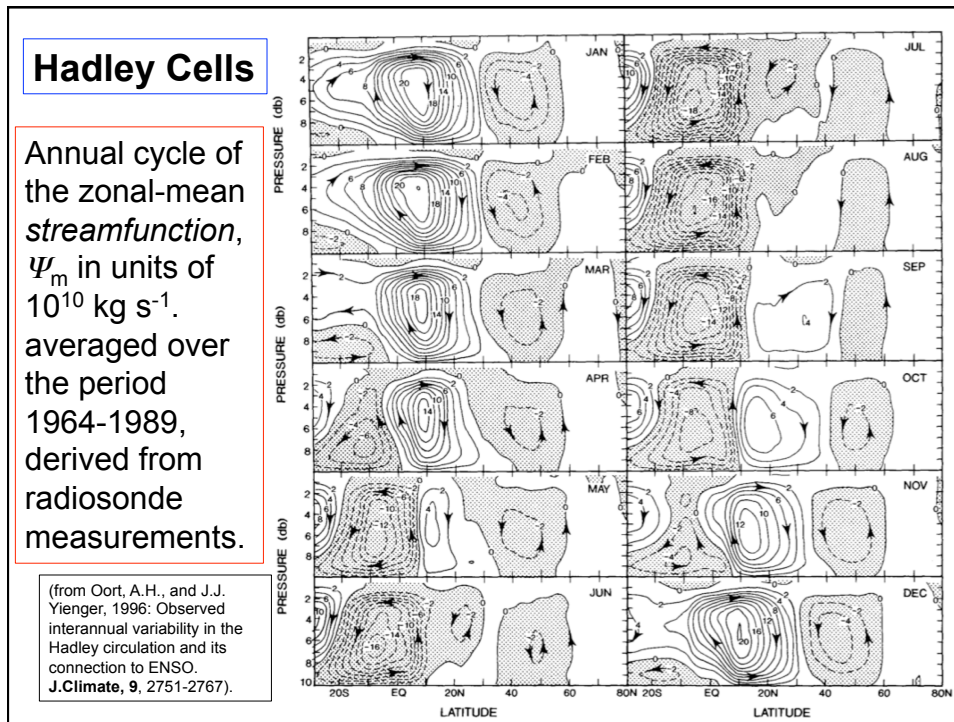
Position of extratropical tropopause

Separation between subtropical jet and stratospheric jet

Zonal mean zonal stratospheric wind reversal

Seasonal cycle of the Hadley circulation





## Important processes

**Radiation**

**Circulation**

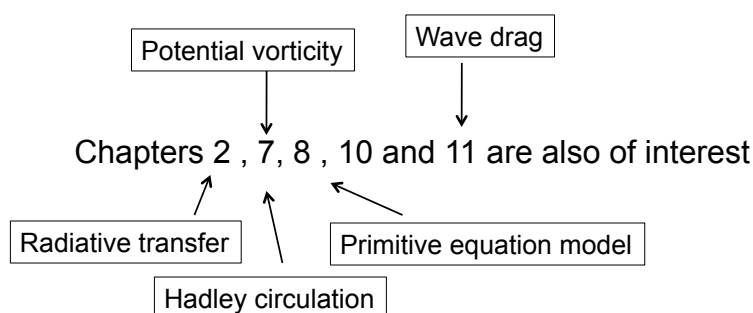
**Energy sources/sinks associated with phase changes of water**

**“Wave drag”**

## Lecture notes

### Chapter 12 of lecture notes on Atmospheric Dynamics

<http://www.staff.science.uu.nl/~delde102/AtmosphericDynamics.htm>



## Program of the C&HC-2

- 23 April: Introduction to radiative transfer; “grey gas”; radiative equilibrium  
*study sections 2.1-2.4 & boxes 2.1-2.4;*  
*(1) problem 12.1 (response time) (0.5)*
- 30 April: Radiatively determined state; Reanalyses  
*(2) problem 12.2 (radiation at TOA; ERA-Interim) (2.0)*
- 7 May: Radiative-dynamical interaction in a dry atmosphere; GCM’s  
*(3) article for review (yes/no); Topic of presentation (GCM)*
- 14 May: Role of water cycle in the general circulation (the ITCZ)  
*(4) problem 12.3 (check of model assumptions) (2.5)*
- 21 May: Role of wave drag in the general circulation (the surf zone)  
*(5) problem 12.5-12.9 (what-if? thought experiments) (1)*
- 4 June: The Hadley-circulation  
*(6) problem 12.12 (Hadley-circulation theory) (1)*
- 11 June: Isentropic coordinates and potential vorticity (inversion)  
*(7) problem 12.10 (isentropic density profile) (1)*
- 18 June: Zonal mean mass- and potential vorticity budget
- 25 June: *(8) presentations on GCM’s (2.0)*
- No exam

## Content of first lecture

### Introduction to radiative transfer

**Absorption of radiation:** Bouguer-Lambert-Beer Law.

**Radiative energy budget equation:** Schwarzschild's equation.

**Radiative equilibrium:** solution of Schwarzschild's equation.

>>Radiatively determined temperature as a function height

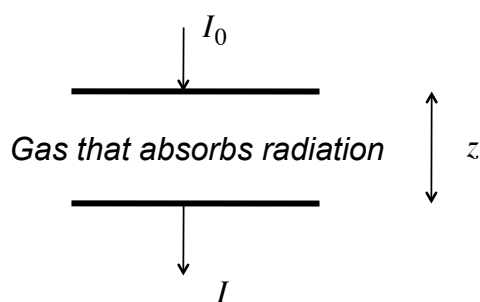
Division of the atmosphere in **troposphere and stratosphere.**

Box 2.2, 2.3 and 2.4 of the lecture notes

<http://www.staff.science.uu.nl/~delde102/C&HC.htm>

### BOX 2.2

## Absorption of radiation



**Bouguer-Lambert-Beer Law:**  $I = I_0 \exp(-\kappa z) \rightarrow dI = -\kappa I dz$

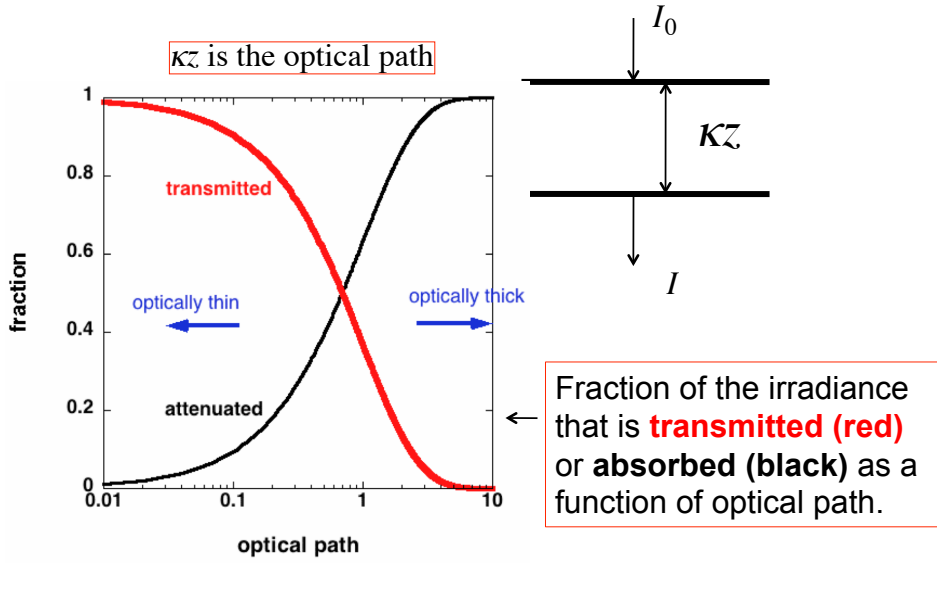
$\kappa$  is the absorption coefficient

$\kappa z$  is the optical path

$\frac{1}{\kappa}$  is the absorption length (unit : m)

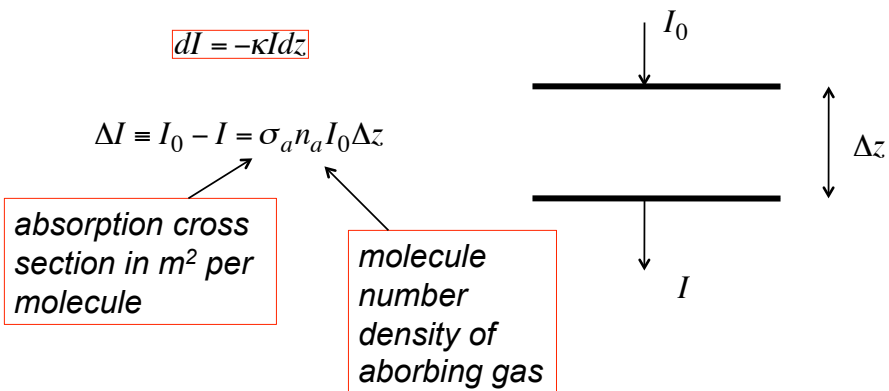
**BOX 2.2**

**Absorption of radiation**



**BOX 2.3**

**Absorption cross section**



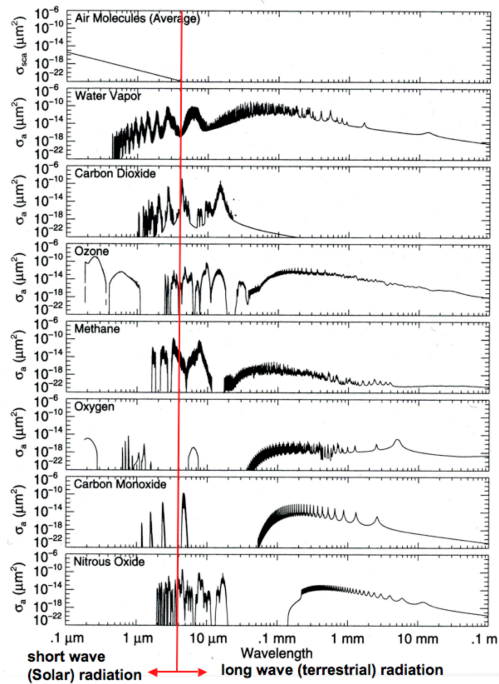


# Absorption cross section

FIGURE 2.48.

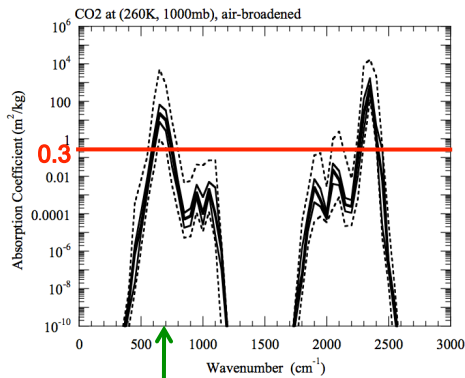
*Absorption cross sections of the strongest absorbing atmospheric constituents*

**Grey gas:** absorption cross section is constant in long wave domain.



# Absorption cross section CO<sub>2</sub>

Absorption cross section CO<sub>2</sub> depends on wavelength



peak long wave radiation spectrum

## **Kirchhoff's law**

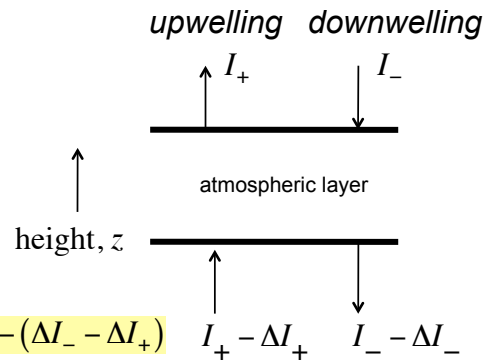
*absorption coefficient = emission coefficient*

Let us

**Investigate the temperature profile  
in an atmosphere which is  
transparent to Solar radiation**

## Upwelling and downwelling radiation

net in from the top =  $I_- - I_+$



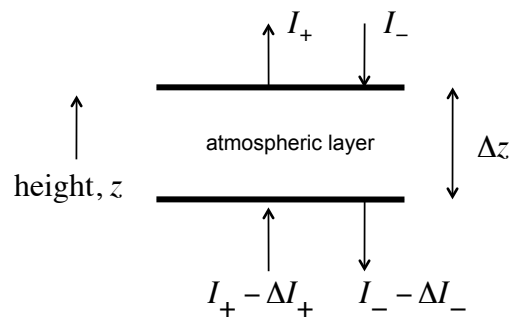
net out from the bottom =  $(I_- - I_+) - (\Delta I_- - \Delta I_+)$

Net absorbed radiation is  $(\Delta I_- - \Delta I_+)$

**Radiation balance:**  $\Delta I_- = \Delta I_+$

### BOX 2.4

## “Planck emission”



Previous slide: **Radiation balance:**  $\Delta I_- = \Delta I_+$

**Fluxes include “Planck emission” of the atmospheric layer!!!**

Black body “Planck emission”:  $B \equiv \sigma T^4$

Atmospheric layer itself emits:  $\kappa \Delta z B$

BOX 2.4

### Schwarzschild's equations

Energy budget:

upwelling radiation:

$$-\Delta I_+ = +\kappa\Delta z(I_+ - \Delta I_+) - \kappa B\Delta z$$

Optically thin layer:

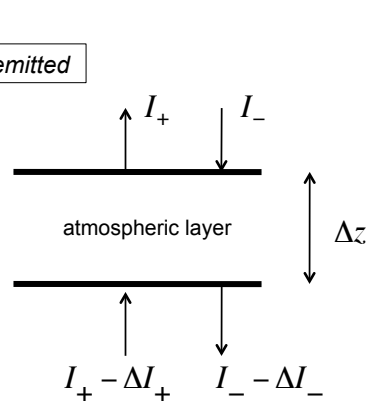
$$-\Delta I_+ = +\kappa\Delta z I_+ - \kappa B\Delta z$$

Likewise:

downwelling radiation:

$$\Delta I_- = \kappa\Delta z I_- - \kappa B\Delta z$$

Previous slide: Radiation balance:  $\Delta I_- = \Delta I_+$



BOX 2.4

### Schwarzschild's equations

Energy budget:

upwelling radiation:

$$-\Delta I_+ = +\kappa\Delta z I_+ - \kappa B\Delta z$$

downwelling radiation:

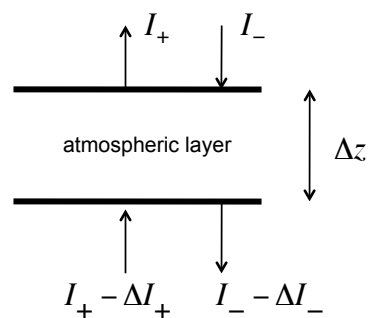
$$\Delta I_- = \kappa\Delta z I_- - \kappa B\Delta z$$

$$\downarrow \quad \Delta I_- = \Delta I_+$$

$$+ \quad I_+ + I_- - 2B = 0$$

and

$$- \quad \Delta I_+ + \Delta I_- = -\kappa(I_+ - I_-)\Delta z$$



**BOX 2.4**

**Continuous homogeneous atmosphere**

**Definition:**  $\bar{I} \equiv I_+ + I_-$  and  $I \equiv I_+ - I_-$

Previous slide:

$I_+ + I_- - 2B = 0$	→	$\bar{I} = 2B$
$\Delta I_+ + \Delta I_- = -\kappa(I_+ - I_-)\Delta z$	→	$\Delta \bar{I} = -\kappa I \Delta z$
$\Delta I_- = \Delta I_+$	→	$\Delta I = 0$

or in general:

	←	$\bar{I} = 2B$
$B \equiv \frac{1}{2} I \delta + B_0$	←	$\frac{d\bar{I}}{d\delta} = I$
	←	$\frac{dI}{d\delta} = 0$

$B_0$  is an integration constant  
Value of  $B$  at  $\delta=0$  (=TOA)

**BOX 2.4**

**Schwarzschild's equation: solution**

$\bar{I} \equiv 2B$	→	$B \equiv \frac{1}{2} I \delta + B_0$	$B_0$ is an integration constant Value of $B$ at $\delta=0$ (TOA)
$\frac{d\bar{I}}{d\delta} \equiv I$	→		
$\frac{dI}{d\delta} = 0$			

**At TOA:**

$I_+ = \frac{1}{4} S_0 (1 - \alpha_p) \equiv Q$

and

$I \equiv I_+ - I_- = I_+ = Q$

$B \equiv \frac{1}{2} (Q \delta + 2B_0)$  for all  $\delta$

$\frac{dI}{d\delta} = 0$  holds for all optical paths

## BOX 2.4

## Schwarzschild's equation: upper bc

$$\bar{I} \equiv 2B \longrightarrow \bar{I} = I + 2I_- \equiv 2B$$

$$\frac{d\bar{I}}{d\delta} \equiv I$$

$$\frac{dI}{d\delta} = 0$$

at TOA this is:  $I = 2B$  or  $Q = 2B_0$

$$B \equiv \sigma T^4 = \frac{1}{2}Q(\delta + 1)$$

$$B \equiv \frac{1}{2}(Q\delta + 2B_0) \quad \text{for all } \delta$$

$$I = Q$$

## BOX 2.4

## Skin temperature

**General solution:**

$$B \equiv \sigma T^4 = \frac{1}{2}Q(\delta + 1) \longrightarrow T = \sqrt[4]{\frac{Q(\delta + 1)}{2\sigma}}$$

$$\text{If } \delta \ll 1 \text{ then } T = \sqrt[4]{\frac{Q}{2\sigma}} \equiv T_{skin}$$

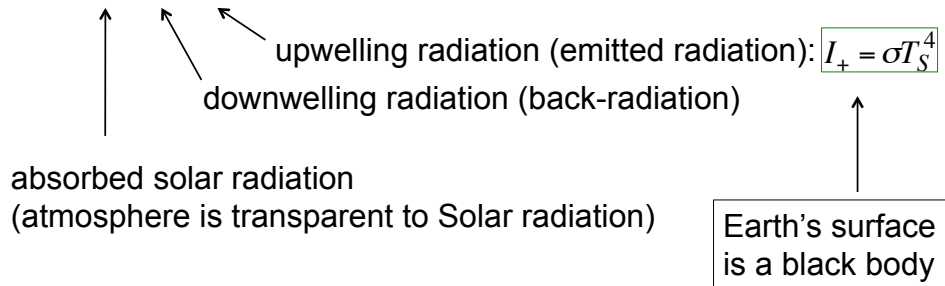
“skin temperature”

$T_{skin}$  is temperature of the **optically thin stratosphere**

**BOX 2.4**

**Radiative equilibrium at surface**

$$Q + I_- - I_+ = 0$$



We had:

$$I \equiv 2B = I_+ + I_- \quad \text{or} \quad I_- = 2B - I_+$$

**BOX 2.4**

**Radiative equilibrium at surface**

So we have:

$$Q + I_- - I_+ = 0$$

$$I_+ = \sigma T_S^4$$

$$I_- = 2B - I_+$$

$$Q + 2B - 2\sigma T_S^4 = 0$$

Surface temperature:

$$T_S = \left( \frac{Q(\delta_S + 2)}{2\sigma} \right)^{1/4}$$

and

optical path at earth's surface

$$B \equiv \sigma T^4 = \frac{1}{2} Q(\delta + 1) = \frac{1}{2} Q(\delta_s + 1)$$

at earth's surface

Atmospheric temperature, just above surface:

$$T_{Sa} = \left( \frac{Q(\delta_S + 1)}{2\sigma} \right)^{1/4}$$

## BOX 2.4

**Radiative equilibrium: full solution**

Atmospheric temperature:

$$T = \sqrt[4]{\frac{Q(\delta+1)}{2\sigma}} = T_{skin}(\delta+1)^{1/4}$$

Surface temperature:

$$T_S = \left(\frac{Q(\delta_S+2)}{2\sigma}\right)^{1/4} = T_{skin}(\delta_S+2)^{1/4}$$

$$T_S > T_{Sa}$$

Atmospheric temperature, just above surface:

$$T_{Sa} = \left(\frac{Q(\delta_S+1)}{2\sigma}\right)^{1/4} = T_{skin}(\delta_S+1)^{1/4}$$

*Surface temperature higher than atmospheric temperature near surface*

## BOX 2.4

**Radiative equilibrium: full solution**

Atmospheric temperature:

$$T = T_{skin}(\delta+1)^{1/4}$$

Surface temperature:

$$T_S = T_{skin}(\delta_S+2)^{1/4}$$

 $\delta$  is related to height or pressure:

$$-\kappa dz = (\sigma_a q_a / g) dp = d(\sigma_a q_a p / g) \equiv d\delta$$

$q$  is the specific concentration by mass of the absorber

$$\delta = \sigma_a q_a p / g \approx \sigma_a r_a p / g$$

$$\delta_s \approx \sigma_a r_a p_s / g$$

$r_a$  is the mixing ratio by mass of the absorber



**BOX 2.4**

**Radiative equilibrium: full solution**

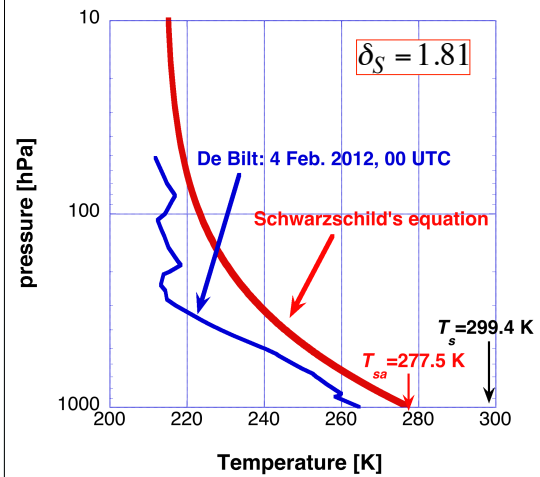
Atmospheric temperature:

$$T = T_{skin}(\delta + 1)^{1/4}$$

Surface temperature:

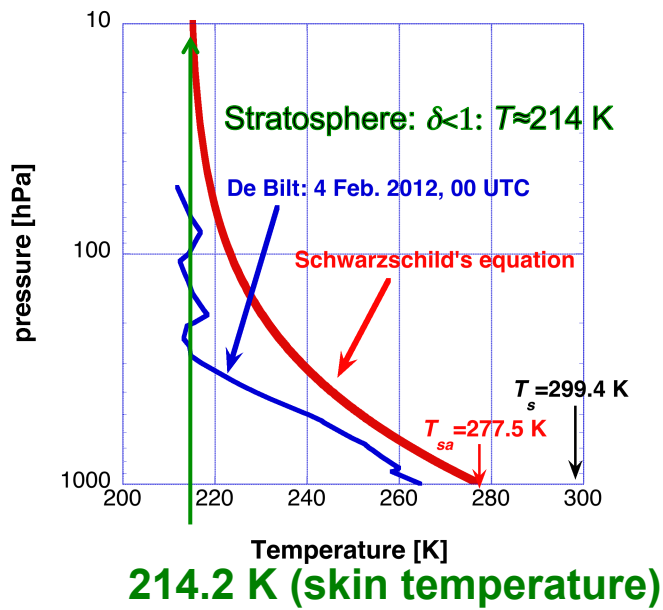
$$T_S = T_{skin}(\delta_S + 2)^{1/4}$$

$\sigma_a = 0.3 \text{ m}^2\text{kg}^{-1}$ ;  $r_a = 390 \text{ ppmv}$   
 $\alpha_p = 0.3$ ;  $S_0 = 1366 \text{ W m}^{-2}$ ;  $p_s = 1000 \text{ hPa}$



**BOX 2.4**

**Radiative equilibrium: full solution**



**Stratosphere is the optically thin upper part of the atmosphere**

## **Assignment 1**

**Problem 12.1 HAND IN ANSWER on 30/4**

(study the theory of box 2.4 and chapter 2.4)

<http://www.staff.science.uu.nl/~delde102/C&HC.htm>

## Next lecture

Wednesday 30/4, 2014, 13:15

*Seasonal cycle of radiative equilibrium and radiatively determined state as a function of latitude, height and time*

*Radiation drives circulation, but circulations also determines radiation: Radiative-dynamical interaction*

<http://www.staff.science.uu.nl/~delde102/C&HC.htm>